

Carbon isotopic variations in species of lichens on an altitude gradient in the South Carpathians and Apuseni Mountains, Romania

Stela Maria Cuna¹, Katalin Bartok², Gabriela Balas¹

¹National Institute of Research and Development for Isotopic and Molecular Technologies,
Donath Street 71-103, P.O. Box 700, Cluj-Napoca 400293, Romania,

e-mail: cuna@oc1.itim-cj.ro

²Babeş-Bolyai University, Kogalniceanu Street 1, Cluj-Napoca, Romania

In the present work, the natural isotopic variations of $^{13}\text{C}/^{12}\text{C}$ in lichens are analyzed in relation to an altitude gradient, to different collection dates, and to different photobiont associations. We have studied the lichens because they are extremely sensitive symbiotic organisms consisting of a fungus (mycobiont) and an algae or cyanobacterium (photobiont) which might react to even slightly polluted air. Lichens are biomonitors with a good accumulation capacity that allow the determination of pollutants deposition in terrestrial ecosystems (Catarino et al., 1991; Nimis et al., 2001).

The lichen species studied were *Cladonia*., *Peltigera canina*, *Hypogymnia physodes* and *Pseudovernia furfuracea*. These lichens were collected in Retezat Mountains, Hășmașu Mare Mountains and Bihor Mountains, Romania. The collection of data was performed during 1978, 1987, 1989, 1994 and 2003. The carbon isotope composition ($\delta^{13}\text{C}$) in organic material from 56 lichen samples has been measured. The lichen samples were combusted to obtain CO_2 for isotopic analysis. CO_2 was purified on cryogenic traps and then analyzed by a dual-inlet isotope ratio mass spectrometer, model ATLAS designed by Varian. The mean standard deviation is $\pm 0.3\%$.

The results show that the $\delta^{13}\text{C}$ values for all studied lichens vary between -21.21% and -26.93% , in the range normally associated with C_3 higher plants (Ehleringer, 1991).

The studied lichens can be separated into two groups on basis of the type of photobiont partner: phycobiont (like *Hypogymnia physodes*) where the photobiont is a green algae, and cyanobiont (like *Peltigera canina*) where the photobiont is a cyanobacteria. We have found a difference between $\delta^{13}\text{C}$ of the two groups: the *Peltigera canina* lichens collected in Bihor Mountains in 1978 showed a mean $\delta^{13}\text{C}$ value of -23.67% , and the *Hypogymnia physodes* collected in the same site and the same date showed a mean $\delta^{13}\text{C}$ value of -22.71% . The difference can be explained by the differences in photosynthetic rates (Maguas and Brugnoli,

1996). The lichens possessing different photobiont associations show different water requirements for the activation of photosynthesis. The cyanobiont lichens require liquid water for photosynthetic CO₂ assimilation, while the phycobiont associations can fix CO₂ in the presence of water vapor. Such differences are related to variation in CO₂ diffusion resistance from air to chloroplasts.

We found a variation of the $\delta^{13}\text{C}$ values in lichens with altitude. For *Hypogymnia physodes* lichens this gradient was from -21.21‰ at 1130m to -26.09‰ at 1520m. We have supposed that the lichens along an altitude gradient are exposed to numerous natural and man-made stress factors and react sensitively to any pollution. The stress factors have a direct or indirect effect on photosynthetic apparatus of lichens, and as a result on the $\delta^{13}\text{C}$ values.

Also, the $\delta^{13}\text{C}$ values of the lichens that were collected at different dates showed a gradient. The *Cladonia* lichens collected in Big Hasmasu Mountains in 1989 year have $\delta^{13}\text{C} = -21.52\text{‰}$ and the *Cladonia* lichens collected in the same mountains, but in 2003 year, have $\delta^{13}\text{C} = -24.75\text{‰}$. The *Pseudovernia furfuracea* lichens collected in the same mountains in 1989 have $\delta^{13}\text{C} = -23.29\text{‰}$, and the *Pseudovernia furfuracea* lichens collected in 2003 have $\delta^{13}\text{C} = -25.03\text{‰}$. All these lichens were collected at the same altitude. The ¹³C content from organic material of these lichens is depleted in the lichens collected in 2003 year in relation to the lichens collected in 1989.

For both species, a decrease in $\delta^{13}\text{C}$ values was associated with the change in the $\delta^{13}\text{C}$ and CO₂ concentration of the CO₂ atmospheric after a lapse of 14 years. This evidence is relevant for understanding the impact of global change on natural vegetation.

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Autor(en)/Author(s): Cuna Stela Maria, Bartok [Bartók] Katalin, Balas Gabriela

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