

# Facies, palaeogeography and stratigraphy of the lower Miocene Traisen Formation and Wildendürnbach Formation (former “Oncophora Beds”) in the Molasse Zone of Lower Austria

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## KEYWORDS

Molasse Basin; Calcite Minimum Interval; upper Ottnangian; Karpatian; Rzehakia Lake System; stratigraphy

## Abstract

A detailed study of OMV wells throughout the Lower Austrian Molasse Basin demonstrates the existence of a distinct and synchronous upper Ottnangian (lower Miocene) stratigraphic signal, the Calcite Minimum Interval (CMI). It corresponds to the depositional phase of the Rzehakia Lake System. This signal is interpreted to be of chronostratigraphic importance as an expression of palaeoclimate and related sea-level change. It is represented by the brackish Traisen Formation, which crops out south of the Danube. The Traisen Formation correlates with sands and shales in OMV wells to the north, termed Wildendürnbach Formation. However, the CMI underlies a marine unit, the so-called “Oncophora Beds” (also known as Rzehakia Beds) as reported from OMV wells in the north at the border to the Czech Republic. We demonstrate that these former “Oncophora Beds” are younger, i.e. of Karpatian age, than originally assumed. Therefore, these deposits cannot be correlated to the late Ottnangian Traisen Formation. This may solve the problem of contradicting interpretations concerning the depositional environment of both units, which were correlated to each other in the past. As no *Rzehakia* bivalves (formerly *Oncophora*) are described from these former “Oncophora Beds”, we recommend to avoid using the term for these turbiditic sands. Instead, we attribute these deposits to the fully marine Karpatian Laa Formation. These new findings, which are in accordance with published data from the Czech Republic, indicate two (in time and space) independent sedimentary systems and sand deposition centres for the Traisen Formation and the massive sands attributed as “Oncophora Beds” around Wildendürnbach. A late Ottnangian system in the south delivered the material of the Traisen Formation from the Alps and a Karpatian system delivered the clastic material of the massive sands of the Laa Formation from the east.

Detaillierte Untersuchungen an OMV-Bohrungen in der niederösterreichischen Molasse-Zone konnten einen markanten und synchronen stratigrafischen Abschnitt, das Kalzit-Minimum-Intervall (CMI), nachweisen, welches einer lakustrinen Phase, dem „Rzehakia Lake System“, entspricht. Es wird angenommen, dass es sich hierbei um ein synchrones Signal ausgelöst durch Klima- und Meeresspiegelschwankungen handelt. Es schließt die brackische Traisen-Formation, welche südlich der Donau aufgeschlossen ist, mit ein. Die Traisen-Formation korreliert mit Sanden und fein-geschichteten siltigen Tonen der Wildendürnbach Formation, welche weiter nördlich in OMV-Bohrungen angetroffen wurden. Allerdings unterlagert die CMI die marinen sogenannten “Oncophora Schichten” (auch Rzehakia Schichten genannt), welche aus OMV-Bohrungen im Norden an der Grenze zur Tschechischen Republik bekannt sind. Es kann gezeigt werden, dass diese “Oncophora Schichten” dem Karpatium zugeschrieben werden müssen. Demzufolge können die “Traisen-Formation” und die “Oncophora Schichten” einander zeitlich nicht entsprechen, wodurch das Problem widersprüchlicher fazieller Interpretationen dieser Ablagerungen gelöst werden kann. Da innerhalb der “Oncophora Schichten” keine namensgebenden Mollusken der Gattung *Rzehakia* (früher *Oncophora*) beschrieben sind, sollte diese Bezeichnung vermieden werden. Stattdessen sollten diese Ablagerungen der marinen, karpatischen Laa-Formation zugerechnet werden. Diese Erkenntnis sowie publizierte Daten aus dem tschechischen Teil des Molasse Beckens lassen darüber hinaus auf 2 unterschiedliche Einspeisungszentren für klastisches Material schließen. Ein Einspeisungszentrum im Süden des Molasse-Beckens liefert während des späten Ottnangiums das klastische Material der Traisen-Formation aus den Alpen. Ein zweites Einspeisungszentrum im Osten liefert im Karpatium die mächtigen Sande der Laa-Formation.

## 1. Introduction

Strata of the Cenozoic Austrian Molasse Basin, a part of the Northern Alpine Foreland Basin system (e.g. Wagner, 1998; Sikula and Nehyba, 2004; Heckeberg et al., 2010) form a thick sedimentary succession in front of the Alpine Orogen, including classical geosites

such as the type locality for the Ottnangian regional Paratethys stage of the Early Miocene (e.g. Steininger et al., 1983; Roegl et al., 1997; Wessely, 2006; Grunert et al., 2010a; Heckeberg et al., 2010; Gebhardt et al., 2013; Pippèr et al., 2018). In contrast to the well-known and thoroughly investigated Molasse strata of

Upper Austria and Bavaria (e.g. Robinson and Zimmer, 1989; Kowalke and Reichenbacher, 2005; Rupp et al., 2008; Heckeberg et al., 2010; Pippèr and Reichenbacher, 2010; Grunert et al., 2012; Grunert et al., 2013; Liebl et al., 2013; Reichenbacher et al., 2013; Sant et al., 2017; Pippèr et al., 2018), the easternmost Lower Austrian Molasse stratigraphy and facies are still partly under-explored (Wagner, 1998; Roetzel et al., 1999; Krenmayr et al., 2002; Piller et al., 2004; Roetzel et al., 2006; Gebhardt et al., 2013; Wimmer-Frey et al., 2013; Egger et al., 2017; Palzer-Khomenko et al., 2018). For instance, the depositional environment of the lower Miocene, upper Ottnangian (Piller et al., 2004; Piller et al., 2007) Traisen Formation (Gebhardt et al., 2013), predominantly massive mica-rich sands between St. Pölten and Tulln, has been a matter of debate over the past two decades. These deposits were seen as part of the Oncophora or Rzehakia Beds, named after the bivalve *Rzehakia* (Rzehak, 1882; Čtyroký, 1972; Mandić and Ćorić, 2007). It was assumed that these sands were deposited under brackish conditions in temporally isolated basins of the Paratethys during the prominent Burdigalian Bur4 sea-level lowstand (Čtyroký, 1972; Mandić and Ćorić, 2007; Piller et al., 2007; Gebhardt et al., 2013). Harzhauser and Mandić (2008) interpreted these deposits as part of an Early Miocene Rzehakia Lake System. During that time, the Central Paratethys Sea disintegrated into several basins with extensive brackish water lakes.

Hamilton (1997) described what he thought equivalent to the “Oncophora Beds” in petroleum industry (OMV) wells from Lower Austria as fully marine turbiditic deposits. His interpretation was mainly based on samples of OMV drill cores from wells around Wildendürnbach (Austrian/Czech border), situated 50 km north of the outcrops of the Traisen Formation. These sediments were deposited at about 1000 m depth with a thickness of several hundreds of metres and are dominated by deep-water turbiditic and massive sands, as described by Hamilton (1997), Palzer-Khomenko et al. (2018) and internal OMV reports (e.g. Sauer and Kuffner, 1997; Kuffner, 2001; both unpublished). These sands characterise a major sand input centre overlying pelitic sediments. According to similar lithofacies sequences, which means lower Miocene pelitic mica-rich sediments overlain by an interval of turbiditic and massive sands, the formerly called “Oncophora Beds” in the north were correlated to the Traisen Formation (former Oncophora/Rzehakia Beds) outcropping south of the Danube. The discrepancy of a brackish Traisen Formation as part of the Rzehakia Lake System correlated to fully marine “Oncophora Beds” has been a matter of debate for years.

Palzer-Khomenko et al. (2018) investigated several wells throughout the Lower Austrian Molasse Basin (LAMB) and demonstrated a major stratigraphic marker interval, the “Calcite Minimum Interval” (CMI). The CMI was interpreted as a significant lithostratigraphic signal of the Rzehakia Lake System affecting upper Ottnangian sediments. It can be identified by various features such

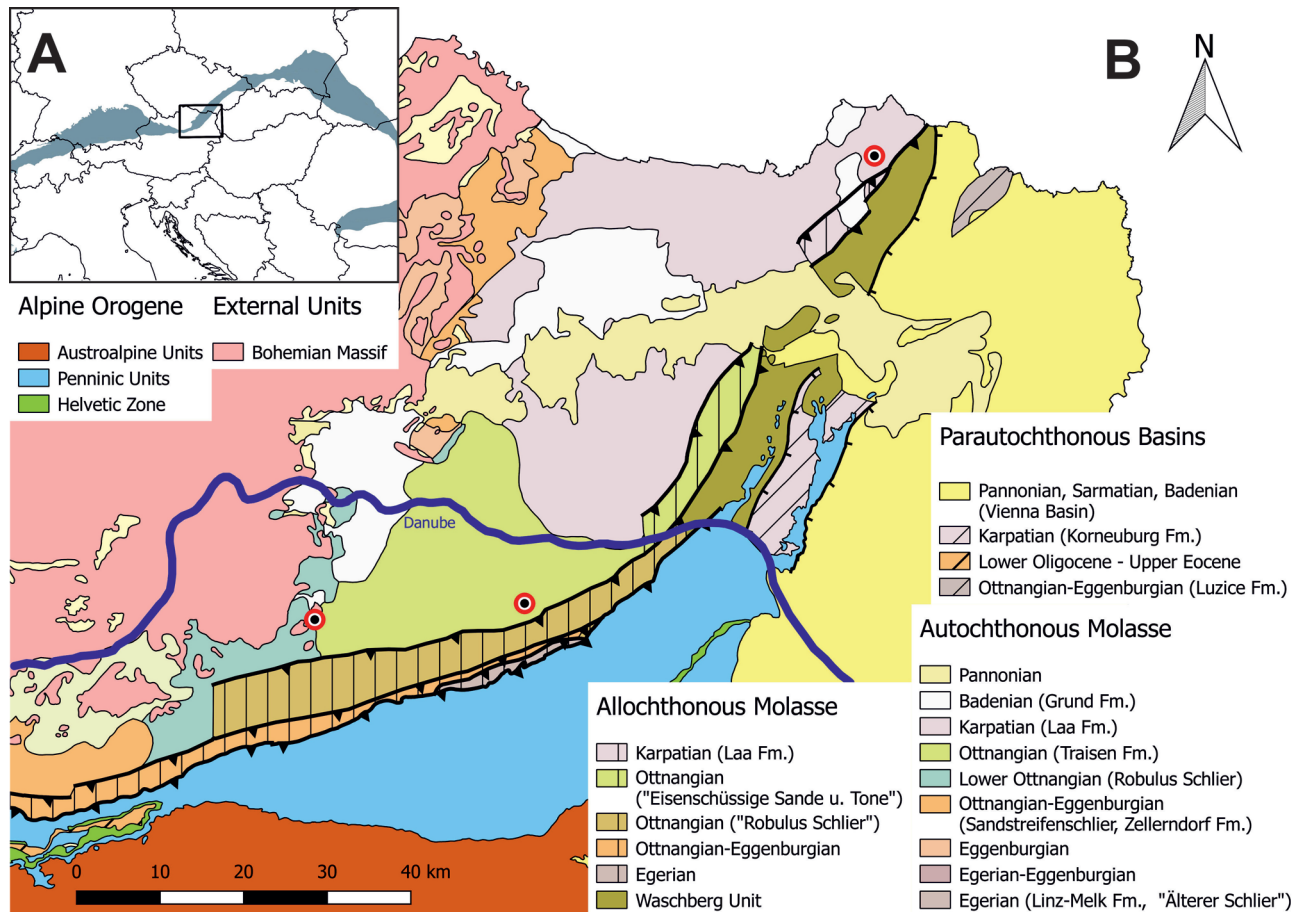
as reduced carbonate contents, increased smectite/illite ratios compared to under- and overlying sediments as well as the absence of (detectable amounts of) pyrite and calcareous nannofossils. For the distal, turbiditic deposits of this interval in the central basinal parts of the LAMB, Palzer-Khomenko et al. (2018) introduced and defined the Wildendürnbach Formation. The present study correlates the Traisen Formation with the Wildendürnbach Formation in the well of Wildendürnbach K4 and demonstrates that the (former) “Oncophora Beds” sensu OMV are overlying this interval and therefore correspond to the Karpatian Laa Formation.

## 2. Geological setting

The LAMB is situated between the Bohemian Massif to the NW, the Rhenodanubian Flysch Zone and the Northern Calcareous Alps (NCA) as part of the Alpine orogen to the S, the Waschberg Unit to the E and the Carpathian orogen to the NE (Fig. 1). To the N, the LAMB continues into the Czech Carpathian Foreland Basin (Nehyba and Sikula, 2007). The sequence of Oligocene-Miocene sediments up to 2000 m thickness is dominated by mica-rich sands and pelites transported from the Alps and the Carpathians. Mica-poor and quartz-rich sands derived from the Bohemian Massif are of subordinate importance and mostly belong to the (older) Linz-Melk Formation (Fig. 1B; Roetzel and Kurzweil, 1986; Krenmayr and Roetzel, 2000). The mica-rich sands and pelites of the lower Ottnangian “Robulus Schlier”, the upper Ottnangian Traisen Formation and the Karpatian Laa Formation make up the major part of the sedimentary infill of the LAMB (Roegl et al., 1997; Krenmayr et al., 2002; Piller et al., 2004; Wessely, 2006; Gebhardt et al., 2013; Wimmer-Frey et al., 2013; Egger et al., 2017; Palzer-Khomenko et al., 2018). Characteristic for these sediments are the garnet-dominated heavy mineral spectra and the fluctuating sand/pelite ratio (Knierzinger et al., 2018). The high frequency, millimetre-to-centimetre intercalations of sands and pelites are termed “Schlier”, which is still in use for many (informal) lithostratigraphic terms in the Austrian Molasse (e.g. “Sandstreifenschlier”, “Robulus Schlier” and “Haller Schlier”; Piller et al., 2004; Wessely, 2006; Rupp et al., 2008; Wimmer-Frey et al., 2013; Egger et al., 2017).

The Traisen Formation is exposed south of the Danube, at the southern to south-eastern rim of the LAMB between St. Pölten and Tulln. According to the definition of Gebhardt et al. (2013), it comprises both a dominating facies of massive sands and a fine-grained lithofacies of subordinated importance. Both sands and pelites show strongly reduced calcite contents (already noticed in the field by using diluted 5% hydrochloric acid) compared to the overlying Laa Formation and the underlying Robulus Schlier. The pelites lack calcareous nannofossils. The only exceptions to the overall carbonate paucity are frequently occurring spherical carbonate concretions with a diameter of up to 1 m.

The turbiditic, time-equivalent continuation of the Traisen Formation to the NE is represented by the Wildendürnbach Formation (Palzer-Khomenko et al.,



**Figure 1.** Adapted maps of Palzer-Khomenko et al. (2018). A: Schematic overview of the Alpine-Carpathian foreland basins and the position of the study area. B: Schematic geological map of the study area and investigated wells (map compilation, Wessely, 2006; Schuster et al., 2015).

2018), mostly known from OMV wells. Its lower and upper lithostratigraphic boundaries were defined by the limits of the CMI. The Wildendürnbach Formation consists of non-calcareous, microfossil-free pelites interrupted by coarse clastic slumps, debris flows and some turbidites. The Wildendürnbach Formation, the Traisen Formation and the locally occurring gravelly Dietersdorf Formation form the Pixendorf Group (Gebhardt et al., 2013).

The upper Ottangian Pixendorf Group overlies the lower Ottangian "Robulus Schlier". The "Robulus Schlier" consists of mica-rich pelitic sediments with subordinate sand intercalations and contains a calcareous nannofossil assemblage indicating the nannofossil zones NN3–NN4 (Čorić and Roegl, 2004; Wimmer-Frey et al., 2013; Egger et al., 2017) of Martini (1971).

North of the Danube, the Pixendorf Group is overlain by the Karpatian Laa Formation. The Laa Formation consists of mica-rich sands and pelites with a largely similar composition compared to the Traisen Formation, except for a higher calcite content and the presence of a Karpatian calcareous nannofossil and foraminiferal assemblage (Čorić and Roegl, 2004; Gebhardt et al., 2013).

Palaeogeographically, a connection between the Central Paratethys and the Mediterranean existed during the early Ottangian (ca. 18 Ma; Krijgsman and Piller, 2012) via Lower and Upper Austria, Bavaria, Switzerland and France. This connection was closed in the late Ottangian

(Roegl and Steininger, 1983; Bieg, 2005). The Upper Austrian and Bavarian parts of the Molasse Basin were isolated and separated from the Lower Austrian part by the structural high of the Amstetten Swell at the southern tip of the Bohemian Massif. The palaeogeographic situation and depositional setting in the LAMB is still under debate (Kapounek et al., 1965; Bieg, 2005; Peña, 2007; Kováč et al., 2017; Sant et al., 2017). The model of an isolated brackish basin belonging to the Rzehakia Lake System (Harzhauser and Mandic, 2008) is supported by mollusc assemblages of the Traisen Formation with the occurrence of the brackish to fresh-water bivalve *Rzehakia* (Mandic and Čorić, 2007; Gebhardt et al., 2013). On the other hand, geologists of OMV brought forward the idea of a fully marine late Ottangian Molasse Basin based on microfossil investigations of well data, which were interpreted as evidence for a prevailing marine connection to the E during the "late" Ottangian (Hamilton, 1997). In contrast, Czech workers already correlated these units to Karpatian deposits in the Czech Republic (Aniwandter et al., 1990; Sikula and Nehyba, 2004; Nehyba and Sikula, 2007).

### 3. Materials and methods

Throughout this study, a total of 280 drill core samples and additional 164 samples of cuttings of seven wells were analysed. Three wells were chosen for the

comparison and correlation of the Traisen Formation and the “Oncophora Beds” sensu OMV around Wildendürnbach: Schaubing (water well, data courtesy by Austrian Geological Survey), Streithofen 1 and Wildendürnbach K4 (Fig. 1). While the shallow (140 m) and marginal well of Schaubing offered a continuous core, the other industrial wells from more basinal and northern parts provide only sparse core material with stratigraphic gaps of 50 m in between. Hereinafter, we present data of all 171 drill core samples of the wells of Schaubing, Streithofen 1 and Wildendürnbach K4.

All samples were powdered and analysed by X-ray diffraction (XRD) and X-ray fluorescence (XRF), as well as for their carbonate content. Thirty samples were chosen for thin sections, 44 for chemical analyses, 46 for organic-walled dinoflagellate cysts investigations (for sample preparation method, see Soliman, 2012) and 67 for calcareous nannofossil investigations by using a modified settling slide method (for sample preparation method, see Burnett, 1998). Tests for washing microfossils were done on 14 samples, but all proved barren of any foraminifera or ostracods. The carbonate content was measured by using a carbonate bomb (Mueller and Gastner, 1971). The bulk mineralogy was analysed by XRD on oriented powdered samples using a Panalytical X'Pert PRO (CuK $\alpha$  radiation, 40 kV, 40 mA, step size 0.0167, 5 s per step) diffractometer. Semi-quantitative information was extracted by using peak intensities as counts per second (cps) of characteristic peaks with low or without overlapping problems with other (present) minerals. Background values were subtracted from peak heights. Dry powders were analysed using a Bruker AXS TRACER

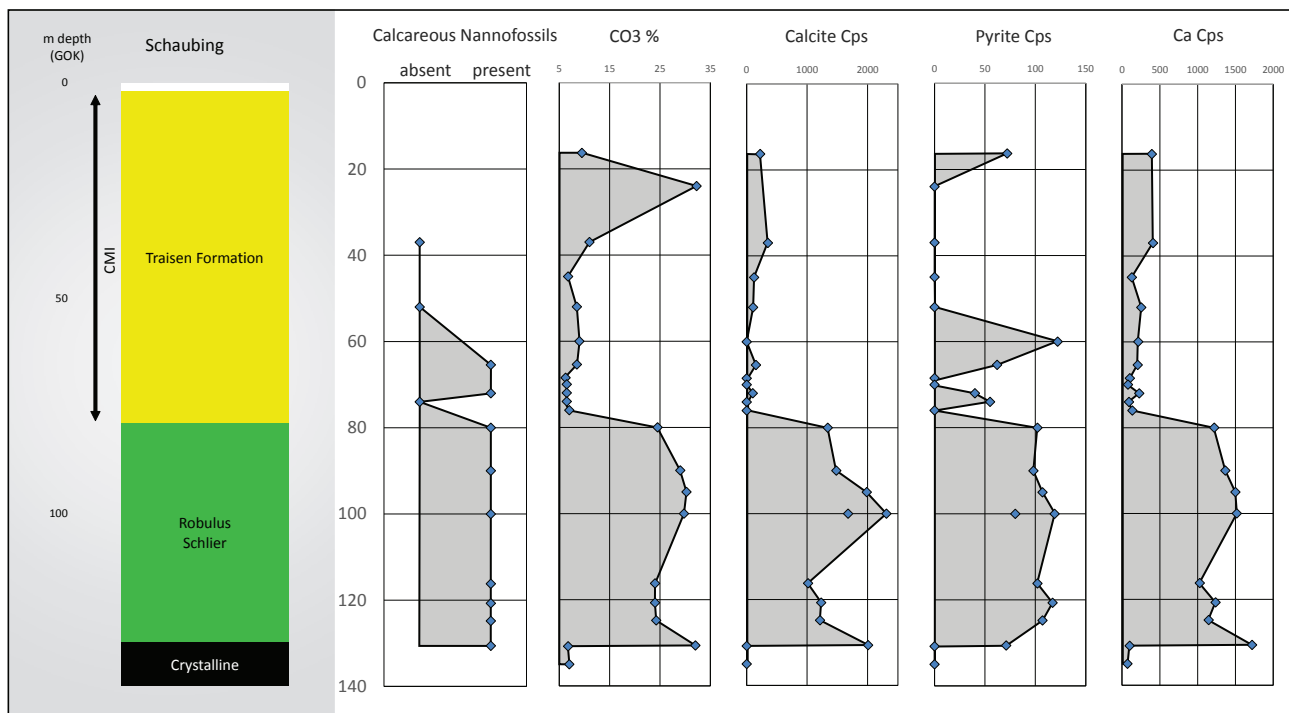
IV-SD handheld ED-XRF instrument (XRF). The measured peak intensities were used for inferring semi-quantitative trends calibrated to internal quantitative data, for which elements/oxides were determined by inductively coupled plasma-mass spectrometry (ICP-MS) by the AcmeLabs (now Bureau Veritas) in Vancouver, British Columbia, Canada.

The data were plotted against depth using carbonate contents (weight%), XRD peak heights (counts per second) characteristic for calcite (3.03 Å) and pyrite (2.70 Å), as well as the XRF counts per second for Ca. Samples of calcite-cemented sands, gravels and concretions were excluded from the profile as they would produce outliers obscuring the CMI signal. They are a common feature within the Traisen Formation and typical for the CMI interval (Gebhardt et al., 2013; Palzer-Khomenko et al., 2018).

## 4. Results

### 4.1. Well of Schaubing

The well of Schaubing offers a 144-m-long continuous drill core from the surface to the crystalline basement of the Bohemian Massif at 130 m depth (Fig. 1). A description of the drill core is also given in Ćorić et al. (2013). The basal granulites are overlain by a conglomerate and monotonous carbonate-rich silty shales. Layers of coarse-grained sands composed of reworked granulites are intercalated up to 112 m. From 112 m to 72 m depth, marly, silty shales containing numerous fish scales become frequent. At 72 m, a first clastic input of fine-grained sands occurs, followed by a clay layer with gypsum rose crystals growing on the drill core. The gypsum is most probably of



**Figure 2.** Adapted depth profiles (GOK = depth below surface) of Schaubing of Palzer-Khomenko et al. (2018). Shown are the counted individuals of calcareous nannofossils, weight% of carbonate and characteristic XRD peak heights (in counts per second) of calcite and pyrite, as well as XRF raw data (counts per second). Please note that the depth scales differ in Figures 2, 4, 5, 6.



secondary origin due to the decay of pyrite. Above 72 m, calcite is not detectable by the hydrochloric acid test. The proportion of sand increases. Above 40 m, massive sands and carbonate-cemented sandstones occur.

Twelve samples from the Schaubing borehole were investigated for calcareous nannofossils and dinoflagellate cysts. Concerning the calcareous nannofossils, the samples between 37 m and 74 m are barren, while the samples belonging to the interval between 80 m and 130.6 m contain moderate-to-poorly preserved specimens (Fig. 2). The most abundant species and taxonomical groups are *Coccolithus pelagicus*, *Reticulofenestra* spp., *Helicosphaera* spp. and reworked nannofossils (for detailed results, see Palzer-Khomenko et al., 2018). Dinoflagellate cysts (Fig. 3) did not occur above 52 m, and only few taxa were found

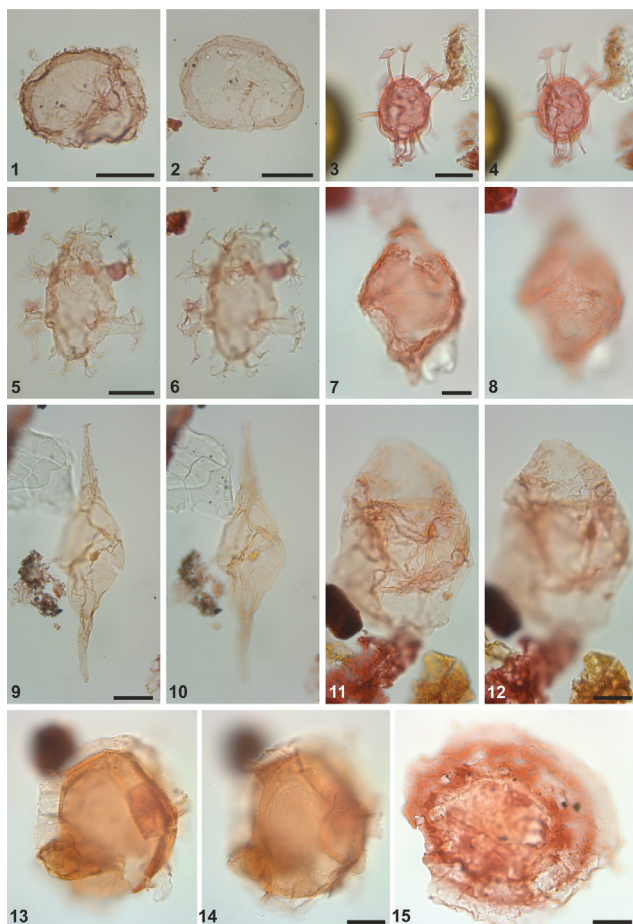
between 76 m and 65.4 m. Most dinoflagellate cysts occurred below 80 m. *Achomosphaera/Spiniferites* spp., *Cleistosphaeridium placacanthum/ancyreum* complex, *Operculodinium centrocarpum*, *Polysphaeridium zoharyi*, *Reticulatosphaera actinocoronata* and *Lingulodinium machaerophorum* were most abundant (Table 1A). *Prasinophytes* spp. were abundantly recorded between 80 m and 100 m.

The carbonate contents (Fig. 2) are higher than 20% for all samples below 80 m depth (except samples of the basement). At 80 m, the carbonate content decreases to <10% in all samples, except in a calcite-cemented sand sample at 24 m. This correlates consequentially with a drop in the calcite XRD peak heights and Ca XRF counts. The XRD peak heights of pyrite are >40 cps at depths below 80 m and decrease to <40 cps between 80 m and 60 m, where strong fluctuations were detected. Except one outlier, pyrite was not detected above 60 m depth.

#### 4.2. Well of Streithofen 1

The well of Streithofen 1 (Fig. 4) is situated in the outcrop area of the Traisen Formation in the southern part of the LAMB (Fig. 1) and provides well-preserved drill cores of 5 m every 50–150 m. The crystalline basement at 1230 m is overlain by fluvialite quartz-rich sand/sandstones and coaly clay/mudstones to siltstones. The coaly beds are tectonically fractured and show pieces of polished slickensides. The facies changes from fluvialite to marine with typical strongly bioturbated bluish sandstones between 1140 m and 1121 m. At 1100 m follows the pelitic “Schlierer” succession of millimetre-to centimetre interbedded silty to clayey shales, yellowish to brownish sands and silts as well as light grey sandstones (Robulus Schlier). Above this, the facies changes to sand-dominated deposits between 850 m and 800 m. A prominent decrease in carbonate content can be recognised. The sediments above 800 m comprise the typical Traisen Formation characterised by low carbonate and high mica contents. The Traisen Formation is continuous up to the surface.

Twenty-nine samples from the Streithofen 1 borehole were investigated for calcareous nannofossils. All samples below 1120 m and above 850 m are barren, except for a few individuals of *Sphenolithus* sp. at 654.8 m and reworked *Sphenolithus ciperoensis* at 151.5 m (Fig. 4). The interval between 1120 m and 850 m contains only very few calcareous nannofossils. The most abundant species is *Coccolithus pelagicus*, followed by scarce *Reticulofenestra* *hagii*, *Sphenolithus moriformis*, *Cyclicargolithus floridanus* and *R. pseudoumbilicus* (for detailed results, see Palzer-Khomenko et al., 2018). Seventeen samples were investigated for dinoflagellate cysts. All samples between 250 m and 990 m are barren. Few individuals of *Cribroperidinium* sp. and undetermined cysts are found in three samples above 250 m. Rare and poorly preserved dinoflagellate cysts occur below 990 m. *Glaphyrocysta* spp., *Cleistosphaeridium* spp., *Cordosphaeridium cantharellus*, *Spiniferites/Achomosphaera* spp. and *Lingulodinium machaerophorum* are the most abundant species (Tab. 1B).



**Figure 3.** Photomicrographs of dinoflagellate cysts (bright field). Scale bar is 20 nm for all figures. 1. *Selenopemphix brevispinosa* Head et al., 1989; Sample WDK4-854.2, Slide1, England Finder E50. 2. *Selenopemphix nephroides* Benedek emend. Bujak et al., 1980; Samples WDK4-725, Slide 1, England Finder G40-3. 3–4. *Melitasphaeridium choanophorum* (Deflandre and Cookson) Harland and Hill, 1979; Sample WDK4-854.2, Slide 1, England Finder P50-2. 5–6. *Distatodinium paradoxum* (Brosius) Eaton, 1976; Sample WDK4-854.2, Slide 1, England Finder S55. 7–8. *Apteodinium spiridoides* Benedek, 1972; Sample WDK4-1254.2, Slide 1, England Finder A-L44. 9–10. *Palaeocystodinium powellii* Strauss et al., 2001; Sample WDK4-854.2, Slide 1, England Finder E50-1. 11–12. *Hystrichosphaeropsis obscura* Habib, 1972; Sample WDK4-1254.2, Slide 1, England Finder T60-3. 13–14. *Pentadinium laticinctum* Gerlach, 1961 emend. Benedek et al., 1982; Sample WDK4-1254.2, Slide 1, England Finder P58-4. 15. *Tuberculodinium vancampoe* (Rossignol) Wall, 1967; Sample WDK4-854.2, Slide 1, England Finder E50-4.

80

Age		Sample	Dinoflagellate cysts																							
			<i>Glaphyrocysta</i> spp.	<i>Deflandrea phosphoritica</i>	<i>Cleistosphaeridium</i> spp.	<i>Cordosphaeridium cantharellus</i>	<i>Dapsilidinium pseudocolligerum</i>	<i>Polysphaeridium zoharyi</i>	<i>Homotryblum</i> spp.	<i>Spiniferites/Achomosphaera</i> spp.	<i>Exochosphaeridium insigne</i>	<i>Hystriocholpoma</i> spp.	<i>Melitasphaeridium choanophorum</i>	<i>Lingulodinium machaerophorum</i>	<i>Lejeunecysta diversiforma</i>	Round Brown cysts	<i>Tuberculodinium vancampoeae</i>	<i>Lejeunecysta</i> spp.	<i>Selenopemphix nephroides</i>	<i>Trinovantedinium</i> sp.	<i>Apteodinium spiridoides</i>	<i>Cribroperidinium</i> sp.	Indet. cysts	Total dinocysts		
Lower Miocene		2014-Strth1 53,8																							3	
		2014-Strth1 151,5	1																					1	5	7
		2014-Strth1 249,3																							2	2
		2014-Strth1 351,2																								
		2014-Strth1 451,9																								
		2014-Strth1 554,8																								
		2014-Strth1 654,8																								
		2014-Strth1 754,5																								
		2014-Strth1 801																								
		2014-Strth1 854,9																								
Ottningian		2014-Strth1 964,8																								
		2014-Strth1 990,5			11	1		1		2			8			5	1				1				30	
		2014-Strth1 1022,6											3	1					2						6	
		2014-Strth1 1084,8	3		8					6			1	5	1	6	2	3	1	1					37	
		2014-Strth1 1119,7	5		8	1				2	6	1													23	
Pre-Eggenburgian		2014-Strth1 1142	10	1	6	5	2	3	18	5														50		
		2014-Strth1 1178,9	3	1	1	3																		8		

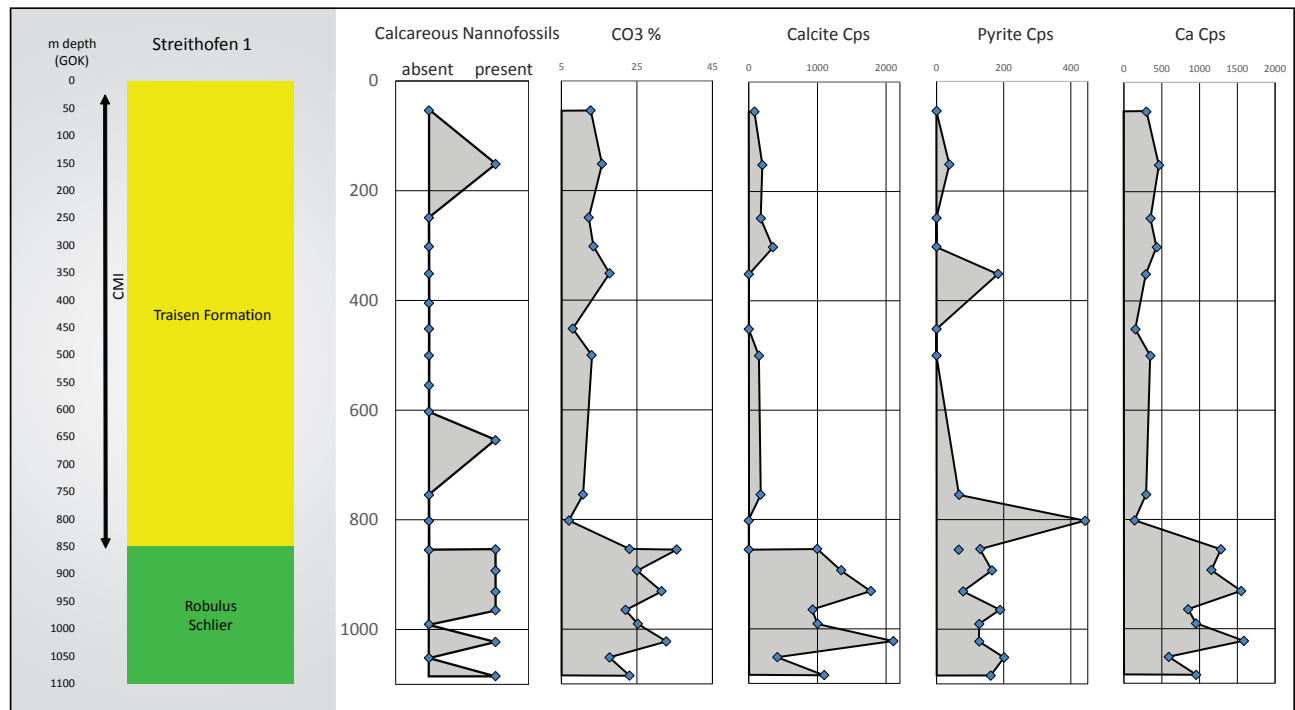
Table 1B: Continued.

Table 1B: Continued.

Table 1C: Continued.

Age	Sample	Lower Miocene	
		Ottangian	
Dinoflagellate cysts	WDK4-725	50	44
	WDK4-854.2	44	23
	WDK4-905		
	WDK4-954.8	25	300
	WDK4-1051		
	WDK4-1054.2	27	2
	WDK4-1095.4	4	1
	WDK4-1154.6	5	
	WDK4-1201.5	87	16
	WDK4-1254.2	53	55
	WDK4-1301		
	WDK4-1353		
	WDK4-1402.6		
	WDK4-1504.5		
	WDK4-1533		
	WDK4-1563	7	5
	<i>Achomosphaera/Spiniferites</i> spp.		
	<i>C. placacanthum/ancyreum</i> complex		
	<i>Selenopemphix nephroides</i>		
	<i>Lejeunecysta globosa</i>		
	<i>Lingulodinium machaerophorum</i>		
	<i>Trinovantedinium</i> sp.		
	<i>Selenopemphix brevispinosa</i>		
	<i>Distatodinium paradoxum</i>		
	<i>Hystrichokolpoma denticulatum</i>		
	<i>Reticulatosphaera actinocoronata</i>		
	<i>Palaeocystodinium powellii</i>		
	<i>Impagidinium</i> sp.		
	<i>Hystrichokolpoma rigaudiae</i>		
	<i>Dapsilidinium pastielsii</i>		
	<i>Spiniferites paradoxum</i>		
	<i>Melitasphaeridium choanophorum</i>		
	<i>Operculodinium centrocarpum</i>		
	<i>Hystrichosphaeropsis obscura</i>		
	<i>Lejeunecysta</i> spp.		
Other Palynomorphs	<i>Polysphaeridium zoharyi</i>		
	<i>Cordosphaeridium cantharellus</i>		
	<i>Batiacasphaera sphaerica</i>		
	<i>Lejeunecysta communis</i>		
	<i>Palaeocystodinium minor</i>		
	<i>Cribrerodinium giuseppi</i>		
	<i>Exochosphaeridium insigne</i>		
	<i>Pentadinium laticinctum</i>		
	<i>Deflandrea phosphoritica</i>		
	<i>Hystrichokolpoma cinctum</i>		
	<i>Homotryblum tenuispinosum</i>		
	<i>Thalassiphora pelagica</i>		
	<i>Glaphyrocysta</i> spp.		
	<i>Apteodinium spiridoides</i>		
	<i>Operculodinium tiara</i>		
	<i>Hystrichokolpoma</i> spp.		
	<i>Homotryblum</i> spp.		
	indet. cysts		
	<b>Total dinocysts count</b>	108	96
	<i>Tasmanites</i> sp.	45	52
	<i>Pterospermella</i> spp.	1	1
	<i>Pediastrum</i> sp.		
	<i>Paralecaniella indentata</i>		





**Figure 4.** Adapted depth profiles of Streithofen 1 of Palzer-Khomenko et al. (2018). Shown are counted individuals of calcareous nannofossils, weight% of carbonate and characteristic XRD peak heights (in counts per second) of calcite and pyrite, as well as XRF raw data (counts per second). Please note that the depth scales differ in Figures 2, 4, 5, 6.

The carbonate content varies between 5% and 70%, whereas most of the samples do not reach values above 40%, except the calcite-cemented outliers (excluded from Fig. 4). The sands overlying the crystalline basement are characterised by carbonate contents mostly <10%. The bioturbated sandstones at 1120 m and the overlying pelites and intercalated sand layers up to 890 m show contents between 15% and 40%. At 800 m depth, the carbonate content drops below 20% and ranges between 9% and 15%. The calcite XRD peak heights (cps) and XRF counts of Ca correlate to the carbonate content. The XRD peak heights of pyrite show a maximum value at 800 m and a fall at 750 m.

#### 4.3. Well of Wildendürnbach K4

The well of Wildendürnbach K4 to the northeast of the study area (Fig. 1) was investigated in detail between 750 m and 1564 m. According to various internal reports of OMV, notes on Self-Potential Logs (SP-Logs) and the OMV database, the so called “Oncophora Beds” should be present between 1250 m and 750 m or between 1000 m and 800 m. In contrast, Palzer-Khomenko et al. (2018) recently identified and defined a new lithostratigraphical unit, the Wildendürnbach Formation, as the distal, turbiditic coeval equivalent of the proximal Traisen Formation.

Drill cores were available every 50 m. In some cases, <50 cm of the original 2–4 m of core (according to internal OMV archives) were left. The material was unsorted, and it was not possible to reconstruct the orientation in most cases. In some cases, the drill core thickness varied, which indicates that the material was partly mixed up

with the drill cores of other wells. Such material as well as drill gouges were excluded from our interpretation. Despite these qualitative drawbacks, the overall characteristics in a larger context could be reconstructed.

The drill cores comprise shales, well-to-poorly sorted brownish-to-yellowish silt/siltstones and light grey sand/sandstones with different degrees of consolidation. The light grey sands are monotonous, micaceous and mostly rich in carbonate. They contain clay clasts and plant remnants. Fining upward trends, erosional surfaces at the base, planar stratification and ripple structures are common features. The brownish slits show convolute bedding, fining upward trends as well as erosional surfaces and contain a much higher amount of rip-up clasts. The hemipelagic-to-pelagic shales show changing carbonate contents. They are mostly millimetre-laminated and contain shell fragments, fish scales and plant remnants.

The core sections below 1530 m show strongly bioturbated fine-grained and massive sands. Fish scales and plant remnants are common, especially on the bedding planes. Between 1500 m and 1250 m, bluish grey clayey shales dominate, which are seen as pelitic background sedimentation. They are interrupted by light grey sands and brownish silts. Between 1250 m and 1150 m, the amount of light grey sands increases. In some OMV internal descriptions, these sands have been previously attributed to the “Oncophora Beds”. Between 1150 m and 1000 m, the amount of light grey sands and brownish silts is strongly reduced. From 1000 m to 800 m, the sand content dominates and increases upwards. This interval represents the “Oncophora Beds sensu stricto”.

At 750 m, the succession is again dominated by laminated shales. The profile in Figure 5 comprises data of pelitic samples only. They represent the fine-grained background sedimentation. The intercalated turbidites and mudflows tend to obscure the CMI signal.

Twenty-three samples were analysed for the calcareous nannofossil contents. Fifteen samples in general and all except one sample between 1090 m and 1504 m are seen to be barren of nannofossils. The remaining samples contain moderate-to-poorly preserved nannofossils, in which the most abundant species and taxonomical groups are as follows: *Coccolithus pelagicus*, *Reticulofenestra* spp. and reworked nannofossils. Lower amounts of *Helicosphaera* spp., *Sphenolithus* spp. and *Braarudosphaera bigelowii* are present (for detailed results, see Palzer-Khomenko et al., 2018). Sixteen samples were analysed for dinoflagellate cysts (Table 1C). The interval between 1533 m and 1301 m is barren. Above 1255 m, *Achomosphaera/Spiniferites* spp. and *Cleistosphaeridium placacanthum/ancyreum* complex are most abundant. *Prasinophytes alga*, *Tasmanites* and *Pterospermella* are recorded sporadically in most samples and are found abundantly at 725 m.

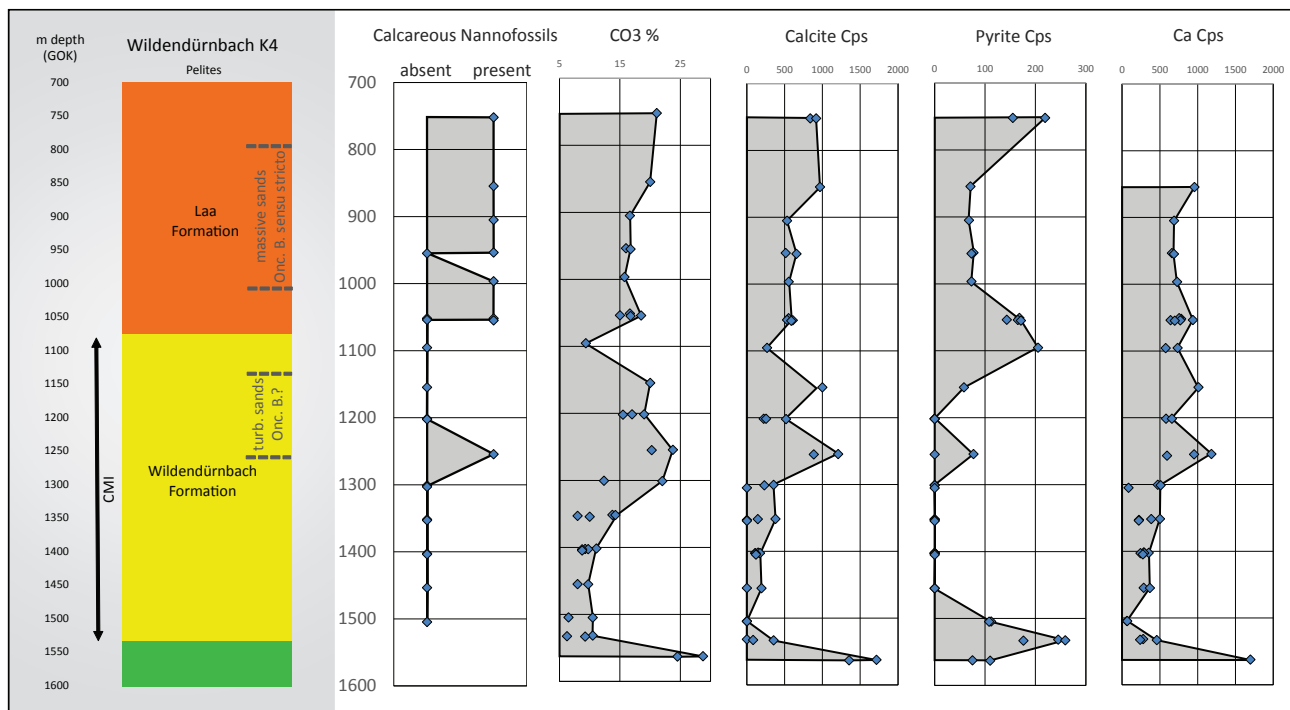
The carbonate contents of the well of Wildendürnbach K4 range between <5% and 30%. The samples at 1560 m show carbonate contents of 25%–30%, followed by a decrease down to 10%. In the following layers, carbonate percentage, calcite XRD counts and Ca XRF counts show large variations within each investigated 5 m drill core, especially for sands. A more continuous profile can be

seen if only pelitic samples are taken into account, which represent the background sedimentation (Fig. 5). This results in the identification of an interval of low calcite values between 1535 m and 1300 m. After a short interval of elevated calcite contents, a sudden decrease occurs again at 1100 m. Pyrite contents decrease after a prominent peak at 1530 m and fully recover at 1100 m.

## 5. Discussion

### 5.1. Biostratigraphy

Following the interpretation of Palzer-Khomenko et al. (2018), the calcareous nannofossil biostratigraphy indicates that the entire section of Schaubing, the investigated interval of Wildendürnbach K4 above 1250 m (as all samples below were barren) and most of the section of Streithofen 1 belong to the interval of NN3 – *Sphenolithus belemnus* Zone (defined at the base by the Last Occurrence [LO] of *Triquetrorhabdulus carinatus* and at the top by the LO of *Sphenolithus belemnus*) to NN4 – *Helicosphaera ampliaperta* Zone (defined by the LO of *Sphenolithus belemnus* at the base and by the LO of *Helicosphaera ampliaperta* at the top) of Martini (1971), which corresponds to the late Early Miocene, i.e. Ottnangian and/or Karpatian age (middle/late Burdigalian, Piller et al., 2007). This age assignment is also proven by dinoflagellate cyst data based on the occurrence of *Cordosphaeridium cantharellus* and *Exochosphaeridium insignis* (Jiménez-Moreno et al., 2006; Soliman et al., 2012). In



**Figure 5.** Adapted depth profiles of Wildendürnbach K4 of Palzer-Khomenko et al. (2018). Shown are counted individuals of calcareous nannofossils, weight% of carbonate and characteristic XRD peak heights (in counts per second) of calcite and pyrite, as well as XRF raw data (counts per second). The extension of the turbiditic sands and massive sands, which were seen as Oncophora Beds in the past (Hamilton, 1997; Sauer and Kuffner, 1997), are indicated by the grey dashed lines. Abbreviations: Turb. = turbiditic; Onc. B. = Oncophora Beds. Note that the depth scales differ in Figures 2, 4, 5, 6.

Streithofen 1, dinoflagellates recovered from two samples at depths of 1142 m and 1178.9 m indicate an age older than Eggenburgian, i.e. early Early Miocene, based on the LO of *Deflandrea phosphoritica* and the co-occurrence of *Cordosphaeridium cantharellus*. It was not possible to derive a higher biostratigraphic resolution than NN2–NN4 for the intervals of interest, which include the late Egerian, Eggenburgian, Ottnangian-Karpatian and early Badenian. The biostratigraphic tools proved to be too inaccurate for a more precise correlation of upper Ottnangian deposits.

## 5.2. Location of the CMI

According to the presented data, the CMI of Palzer-Khomenko et al. (2018) can be identified in the presented sections. A correlation of the CMI for these wells and a comparison to the outdated OMV profile and stratigraphy (Aniwandter et al., 1990) can be seen in Figure 6. The decrease of calcite contents is followed by the decrease of pyrite contents with a delay of 20–100 m. The onset of the CMI is defined by the drop of calcite and the disappearance of calcareous nannofossils (Palzer-Khomenko et al., 2018). Therefore, the CMI starts at 80 m in the section of Schaubing, between 850 m and 800 m in the section of Streithofen 1 and below 1535 m in the section of Wildendürnbach K4. Note that the resolution in Schaubing is much higher due to the continuous drill core, which allows a more accurate localisation of the CMI onset. At Streithofen 1 and Wildendürnbach K4, the gaps between the drill cores cause uncertainties of c. 50 m. At Schaubing and Streithofen 1, the CMI continues up to the surface, where the Traisen Formation crops out.

At Wildendürnbach K4, the localisation of the upper limit of the CMI remains problematic due to the occurrence of detrital and concretionary calcite-bearing sands. If only samples of the fine-grained background sedimentation (Fig. 5) are taken into account, the recovery of carbonate, calcite and Ca can be localised between 1300 m and 1250 m. Calcareous nannofossils occur sporadically in one sample at 1254 m and more continuously from 1054 m upwards. Pyrite XRD peak heights were not detected for all samples between 1450 m and 1350 m and for many (especially for those representing the background sedimentation) samples up to 1150 m. Therefore, the core of the CMI can be identified between 1500 m and 1300 m.

Whether the interval between 1250 m and 1050 m should be seen as part of the CMI or not constitutes a more challenging question. While  $\text{CO}_3$ , calcite and Ca steadily increase between 1300 m and 1250 m, the pyrite contents remain low and calcareous nannofossils remain absent until 1100 m and 1050 m (with one exception of marly silts at 1250 m). A last snap of low  $\text{CO}_3\%$  and calcite values occurs at 1100 m. According to the definition of Palzer-Khomenko et al. (2018), the upper boundary of the CMI is marked by the “sustainable increase of pyrite contents and the coeval re-occurrence of calcareous nannofossils”. These conditions are best fulfilled for the well interval between 1100 m and 1050 m. We therefore infer the upper limit of the CMI, and therefore the lithostratigraphic

boundary between the Wildendürnbach Formation and the Laa Formation, to lie at c. 1075 m in this well.

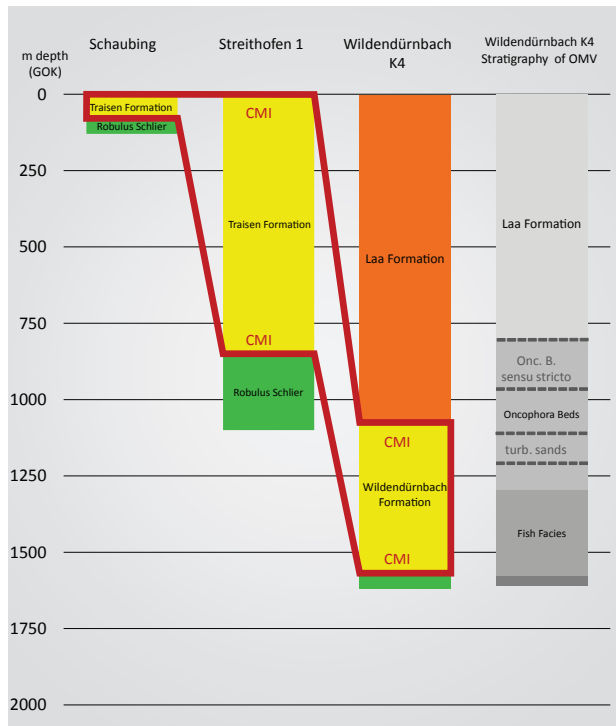
However, the CMI was clearly identified in the interval between 1500 m and 1300 m, which is significantly below the interval of the former “Oncophora Beds sensu stricto” starting at 1000 m. The turbiditic sands around 1200 m fall into the upper part of the CMI but clearly overlie the main part of the CMI. Much of the OMV interpretation of the “Oncophora Beds” as open marine sediments is based on the presence of microfossils. Unfortunately, Hamilton (1997) did not describe the exact origin of the samples on which the interpretation of a marine environment is based. Nevertheless, it can be assumed that the analysed microfossils were derived from the micro- and nannofossil-bearing pelitic beds intercalated between the massive sands between 1000 m and 800 m rather than from the fossil-free interval below. Sedimentary structures typical for turbidites, which were also taken as arguments for marine conditions, are indicators of considerable water depth and sedimentation processes but have no significance concerning the reconstruction of water salinity.

## 5.3. A revised correlation of Traisen Formation and “Oncophora Beds” sensu OMV

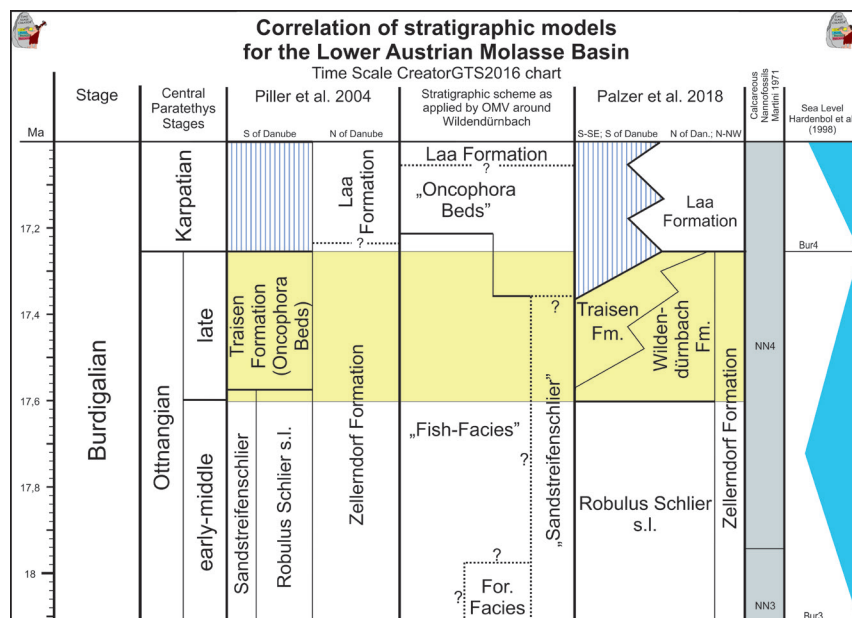
Palzer-Khomenko et al. (2018) inferred a basin-wide synchronous event during the late Ottnangian causing a carbonate crisis and the CMI, which may correspond to the environmental change during the time interval of the brackish Rzehakia Lake System proposed by Harzhauser and Mandic (2008). Consequently, the sediments analysed by Hamilton (1997) attributed to a fully marine facies do not correlate chronologically to the Traisen Formation and to the Rzehakia Lake System comprised by the CMI (Fig. 6). Thus, the so-called “Oncophora Beds” sensu OMV around Wildendürnbach, especially the typical massive sands between 1000 m and 800 m at Wildendürnbach K4, overlie the CMI and are inevitably younger. Following the definition of Piller et al. (2007, p. 155), who described the Ottnangian in the Central Paratethys as a “strictly twofold stage with a normal marine development in its lower part and a predominance of restricted marine to fresh water environments in its upper part”, we interpret the CMI as an excellent basin-wide proxy for the late Ottnangian.

Aniwandter et al. (1990) have previously described the contradicting interpretations for the so-called “Oncophora Beds” at the Czech border. While Austrian workers attributed them to the Ottnangian by correlation with the Traisen Formation further to the south, geologists from the Czech Republic assumed a Karpatian age based on the presence of the Ostracoda *Senesia* aff. *vadaszi* (Aniwandter et al., 1990).

According to these findings, we attribute the former sand-rich “Oncophora Beds” sensu OMV in the area of Wildendürnbach to the Karpatian and thus these form part of the Karpatian Laa Formation (Roegl et al., 1997). The underlying pelitic CMI sediments at Wildendürnbach K4 comprise the Wildendürnbach Formation (sensu Palzer-Khomenko et al., 2018).



**Figure 6.** Correlation of the Calcite Minimum Interval (CMI) for the wells of Schaubing, Streithofen 1 and Wildendürnbach K4. The OMV internal short profile is shown in grey scales for comparison. The upper Ottnangian Wildendürnbach Formation is underlying the "Oncophora Beds" sensu OMV. Please note that the typical massive sands of the "Oncophora Beds" sensu OMV start in Wildendürnbach K4 at 1000 m. The pelites interrupted by sand packages (turbiditic sands) between 1200 m and 1000 m were attributed to the Oncophora Beds in some, but not in all, OMV internal descriptions. Please note that the depth scales differ in Figures 2, 4, 5, 6.



**Figure 7.** Comparison of stratigraphic concepts for the lower Miocene Ottnangian and early Karpatian Lower Austrian Molasse Basin (LAMB) of Piller et al. (2004), of OMV internal reports, of Hamilton (1997) and of Palzer-Khomenko et al. (2018). The Ottnangian/Karpatian boundary correlates to the transgression following the Bur4 lowstand (Hardenbol et al., 1998) and the termination of the Rzehakia Lake System (Harzhauser and Mandic, 2008). Palzer-Khomenko et al. (2018) defined stratigraphic units using the CMI signal. While early, middle and late Ottnangian boundaries are distinguished in Bavaria (Heckeberg et al., 2010; Pippèrr and Reichenbacher, 2017), a corresponding definition and differentiation are unknown in Lower Austria. Abbreviations: For. Facies = foraminiferal facies; Robulus Schlier s. l. = Robulus Schlier sensu lato; Fm. = Formation. Chronostratigraphy based on Time Scale Creator Version 7.1 (<https://engineering.purdue.edu/Stratigraphy/tscreator>) using Gradstein et al. (2012).

The stratigraphic concepts of Piller et al. (2004), OMV reports and descriptions of Palzer-Khomenko et al. (2018) are compared in Figure 7. The yellow box marks the position of the CMI and the Rzehakia Lake System of Harzhauser and Mandic (2008) around the Burdigalian Bur4 sea-level lowstand (Piller et al., 2007). Palzer-Khomenko et al. (2018) propose to define the upper and lower limits of the Traisen Formation and the Wildendürnbach Formation, the upper limit of the Zellerndorf Formation and the Robulus Schlier as well as the lower limit of the Laa Formation by the CMI signal.

#### 5.4. Palaeogeography

Sikula and Nehyba (2004) analysed the numerous drill logs and seismic reflection profiles north of Wildendürnbach and recognised 13 well log facies. The resulting thickness maps indicate a major Karpatian input centre directly north of the Austrian border at the Carpathian fold and thrust front. Comparisons to well logs from Wildendürnbach K4 indicate strong similarities to the interval of the "Oncophora Beds". We therefore correlate the "Oncophora Beds" sensu OMV around Wildendürnbach K4 to the large Karpatian clastic input centre described by Sikula and Nehyba (2004).

Therefore, we assume different provenance areas and sediment input centres for the Traisen Formation in the south and for the former "Oncophora Beds" in the north (Fig. 8). The sediments of the Traisen Formation were delivered by a late Ottnangian river system probably from the Alpine Orogen into the Rzehakia Lake System. Heavy mineral data of Knierzinger (2015) indicate mainly a delivery of material from the Austroalpine Crystalline Complex south of the Rhenodanubian Flysch Zone and the Northern Calcareous Alps (NCA). Therefore, an Enns-type river system overflowing or cutting through the NCA was necessary to deliver the clastic deposits of the Traisen Formation. Following Dunkl et al. (2005) and Frisch et al. (1998), this river system was caught at some point by the NNE-SSW-oriented fault systems related to lateral extrusion and deflected to the east probably to the Vienna Basin following prominent fault systems like the Mur-Mürz fault (Salcher et al., 2012). Kováč et al. (2004) described coeval depocentres at the SW margin of the early Vienna Basin. Since that time, the uprising NCA acted mainly as a watershed south of the LAMB. Sands of the "Oncophora Beds" sensu OMV of Wildendürnbach K4 were delivered by Karpatian-aged pro-delta turbiditic flows from the Carpathian fold and thrust front represented by the initial Waschberg Zone in the east into a



marine LAMB. A delivery from the eastern approximating Waschberg Zone was previously stated by Hamilton (1997). According to the work of Sikula and Nehyba (2004), the activity of the Karpatian sedimentary system and depocentre was lower at the beginning of the Karpatian (Fig. 8C) and increased with time due to the approaching Waschberg thrust units (Fig. 8D).

## 6. Conclusions

The identification of the CMI as a lithological and chemical marker of the Rzehakia Lake System, which represents a chronostratigraphic basin-wide upper Ottnangian key horizon (Palzer-Khomenko et al., 2018) in the well sections of Schaubing, Streithofen 1 and Wildendürnbach K4, allows a more precise correlation between these wells and a new stratigraphic concept in the LAMB. While the Traisen Formation (Gebhardt et al., 2013) in the south is clearly a part of the CMI, the former “Oncophora

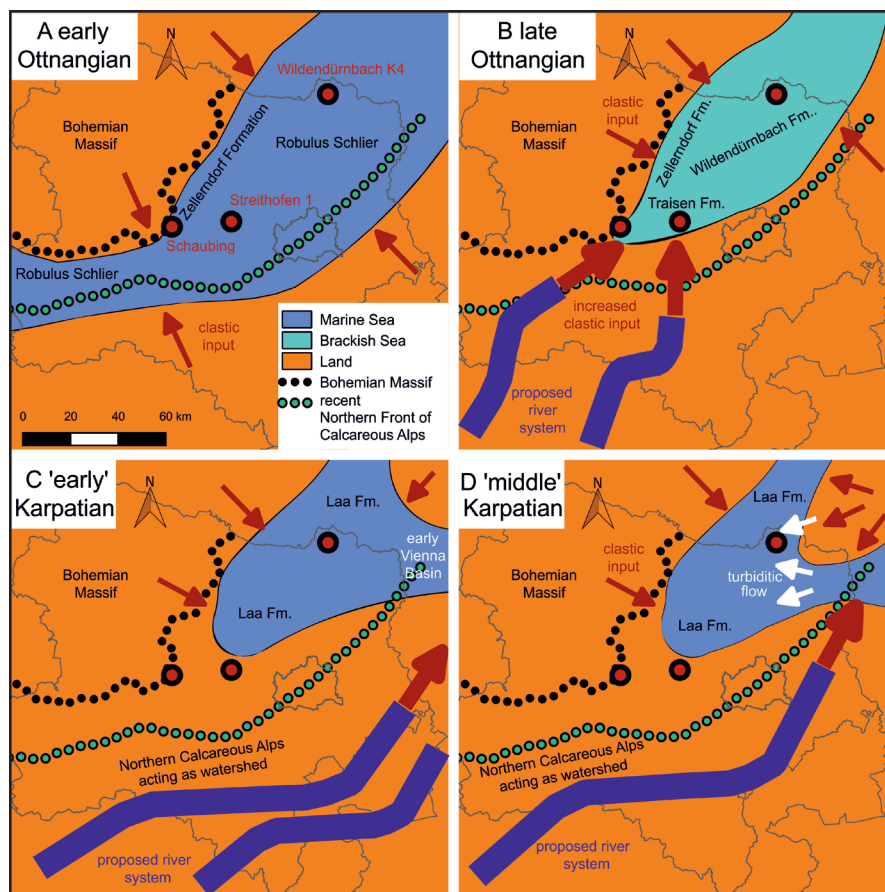
Beds” in the well of Wildendürnbach K4 (Hamilton, 1997) are overlying this interval. Therefore, these “Oncophora Beds” do not represent upper Ottnangian but rather Karpatian deposits. A comparison with data from Sikula and Nehyba (2004) indicates that these sands were delivered from an eastern Karpatian-aged input centre at the Carpathian fold and thrust front. Therefore, we assume two spatially and temporally independent input centres: (1) an older, late Ottnangian input centre to the south delivering the sands of the Traisen Formation into the brackish Rzehakia Lake System during the Bur4 sea-level lowstand at around 17.3 Ma (according to Krijgsman and Piller, 2012), and (2) a younger, Karpatian input centre in the north delivering the sands of the former “Oncophora/Rzehakia Beds” sensu OMV around Wildendürnbach into a fully marine environment after the major Karpatian transgression, i.e. the sea-level rise at the beginning of the highstand Bur4 of the Burdigalian,

younger than 17.2 Ma (Roetzel et al., 2014). However, the timing of the Ottnangian/Karpatian boundary is still under debate, and a considerably younger age of 16.8 Ma was considered by Reichenbacher et al. (2013)

Thus, the problem of the contradicting interpretations concerning depositional environments (shallow-brackish versus deep-water fully marine, Hamilton, 1997; Mandić and Ćorić, 2007; Gebhardt et al., 2013) can be solved as both interpretations apply correctly and as the “Oncophora Beds” in the wells around Wildendürnbach are in fact Karpatian in age. As there is no *Rzehakia* described for these sands, we recommend to avoid the term “Oncophora Beds” and attribute these sediments to be a part of the Laa Formation.

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**Figure 8.** Palaeogeographic evolution of the Lower Austrian Molasse Basin (LAMB) based on reconstructions for the Early Miocene, Ottnangian and early Karpatian boundaries considering reconstructions by Beidinger and Decker (2014), Bieg (2005), Sikula and Nehyba (2004), Frisch et al. (1998), Dunkl et al. (2005) and Kováč et al. (2004). Limits of exposed Bohemian Massif rocks and the recent northern front of the NCA are based on Beidinger and Decker (2014). A. The connection to the western part of the Central Paratethys (west of Amstetten) was a large open seaway during the early Ottnangian (Grunert et al., 2010b). B. This Molasse seaway is closed by tectonic processes driven by northwards moving Alpine thrust systems and the eustatic Bur4 regression during the late Ottnangian. Isolation of the LAMB triggers an environmental crisis, creates the Rzehakia Lake System (Harzhauser and Mandić, 2008) and causes the CMI during the sea-level lowstand (Palzer-Khomenko et al., 2018). In the south, a large clastic input centre develops. C. The “early” Karpatian transgression and the beginning subsidence of the Vienna Basin re-established connections to the east. D. This transgression was followed by the development of a “middle” Karpatian clastic input centre delivering material from the east at the Austrian/Czech border (Sant et al., 2017).

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